

Rocks of the Newark Basin Jd Lowest Lower Jurassic **INTRUSIVE CONTACTS** Newark Supergroup Jurassic Conglomerates of the Jp border fault areas ≥JTicsh Brunswick Jf Lower Jurassic JRps JRpst JRpm UNCONFORMITY Rocks of the Peapack Klippe Lowest Middle and Upper Lower Ordovician UNCONFORMITY Lehigh Valley Sequence Ordovician Beekmantown Group Paleozoic Lower Ordovician Obl Kittatinny Lowest Lower Ordovician and Upper Cambrian Ordovician and Cambrian O€a Supergroup Middle and Lower Cambrian Cambrian Lower Cambrian UNCONFORMITY Rocks of the Reading Prong Byram Intrusive Suite Ybh Yba **INTRUSIVE CONTACTS Metasedimentary Rocks** Yk Yp Yb UNCONFORMITY Precambrian Proterozoic Losee Metamorphic Suite Ylo Ylb **Rocks of Uncertain Origin** Yd | Ya | Yu

CORRELATION OF MAP UNITS

DESCRIPTION OF MAP UNITS

Rocks of the Newark Basin

Intrusive Rock

Diabase - Discordant, stock-like intrusives. Darkgray to greenish-black, coarse- to fine-grained diabase composed mostly of calcic plagioclase and augite (calcic clinopyroxene). Very fine grained where chilled in contact with sedimentary rock. Thickness not known. Elliptical diabase body near Peapack has small remnant of pre-Mesozoic rock on south slope, suggesting nearly vertical ascent of magma through basement

Newark Supergroup

Preakness Basalt - Dark-greenish-gray to black, very fine-grained, very hard, moderately dense rock composed mostly of intergrown calcic plagioclase and pyroxene. Crystals are generally less than 0.04 inch in diameter. In places the basalt contains large feldspar crystals (0.5 inch or larger), or small spherical to tubular cavities (gas-escape vesicles), some filled with zeolite minerals or calcite. Consists of two thick (more than 475 feet each) sequences of basalt flows with thin (10 to 35 feet) beds of siltstone between them (Fedosh and Smoot, 1988). Siltstone beds have not been observed in map area. The basal 65 feet of the lower flow is commonly highly vesicular or brecciated. Radiating fans mostly 15 to 40 feet across consisting of slender, curved basalt columns (8 to 24 inches wide) are common in the upper flows, less common in lower ones. Thickness of formation in map area is about 1050 feet.

Feltville Formation - Mostly fine sandstone and coarse siltstone, brownish-red to light grayishred, moderately well-sorted, commonly crosslaminated; interbedded with brownish-red mudstone, indistinctly laminated; bioturbated, mostly by worm burrows; calcareous in places, with clay and hematite matrix. Sandstone commonly contains more than 15 percent feldspar. A thin bed (0 to 8 feet thick) of black carbonaceous limestone and gray calcareous siltstone occurs near the base. It contains fish and plant fossils, and thermally mature hydrocarbons (Manspeizer and Olsen, 1981). Maximum thickness of formation in map area is about 500 feet.

Orange Mountain Basalt - Dark-greenish-gray basalt composed mostly of calcic plagioclase and augite; crystals smaller than 0.04 inch. Consists of two major flow sequences, each about 265 feet thick (Fedosh and Smoot, 1988) separated in places by a weathered zone. Lower flow is mostly massive, with widely spaced curvilinear joints; upper flow contains one or more zones of slender (6 to 12 inches across) columns.

Maximum thickness in map area is 530 feet. **Passaic Formation**

Sandstone and Siltstone - Medium- and fine-grained sandstone, grayish-red to brownish-red, thick-to very thick-bedded (1.5 to 15 feet thick), with some pebble layers or scattered pebbles (mostly quartzite). Interbedded with brownish- to purplish-red coarse siltstone, indistinctly laminated or ripple cross-laminated, irregularly fissile, and calcareous in places. Some siltstone beds contain nodules, desiccation cracks, or root casts, which are characteristic of fossil soils. Bedding units consist of thick (6 to 15 feet) finingupwards cycles of fluvial channel fill and overbank deposits. Sandstone beds are coarser and thicker in the vicinity of alluvial fan-conglomerate

Siltstone and Mudstone - Coarse-grained siltstone to very fine-grained sandstone, brownish- to purplish-red, thin (0.2 to 1 foot) to medium bedded (1 to 3 feet) with irregular planar, trough, and ripple cross-stratification. Siltstone channelfill deposits are interbedded with silty mudstone beds which are indistinctly bedded, bioturbated, commonly micaceous, and have nodules and calcareous matrix in many places. Maximum

thickness exceeds 1500 feet in map area.

Mudstone - Silty mudstone and argillaceous siltstone, brownish-red, thin-bedded, indistinctly laminated to slightly fissile where micaceous. Mudflat-deposited mudstone beds are soft, crumbly, and commonly calcareous. Lamination is rare in the mudstone, probably because of extensive reworking by organisms during deposition (bioturbation). Channel-fill siltstone beds are mostly 0.2 to 0.5 foot thick, contain have ripple cross-lamination. Mi nimum thickness exceeds 2500 feet in map area.

Conglomerates of the Border Fault Areas Limestone-clast - Pebble to boulder conglomerate with mostly subangular clasts of dolomitic limestone in a matrix of brownish- to purplish-red sandstone to mudstone. Near fault zones the matrix is commonly light-gray to white. Bedding is thick to massive (1.5 to 15 or more feet thick). Clast orientation is mostly indistinct, as in debris flows. Well exposed in abandoned quarries where it was formerly quarried for agricultural lime. Thickness unknown; probably exceeds 1000 feet near the border fault.

Quartzite-clast - Pebble to boulder conglomerate with mostly subrounded quartzite and quartz clasts in matrix of light-red sand to brownish-red silt. Bedding is thick (2 to 6 feet) to very thick (6 to 15 feet). Imbrication and bedding-plane orientation are weak, especially in coarser beds. Quartzite conglomerate occurs adjacent to the border fault northwest of Gladstone, and forms part of a composite alluvial fan sequence south of Pottersville. The conglomerate is coarsest and most thickly bedded adjacent to the border fault, corresponding to the position of alluvial fan apexes. Maximum clast size decreases with distance from the border fault and interbeds of coarse- to fine-grained red sandstone become thicker and more abundant. Maximum thickness probably exceeds 500 feet.

Shale-clast - Pebble conglomerate with green, gray, and purplish-red shale clasts mostly 0.5 to 3 inches long, in a matrix of shale chips and clay. Clasts are mostly flat, elongated, and have rounded edges and corners. Clasts are commonly imbricated at low angles to bedding, dipping in the upstream direction of paleoflow.

Thickness probably exceeds 1000 feet.

Jutland klippe upper unit B (Perissoratis and

Rocks of the Peapack Klippe

others, 1979; Lyttle and Epstein, 1987; Drake and others, 1988) - Heterogeneous sequence of interbedded red and green shale, interbedded fine-grained graywacke siltstone and shale, and yellow, red, green, tan, and gray shale. Much of the rock is an autoclastic melange in a scalycleaved matrix. Some shale beds contain conodonts (Ethington and others, 1958) and sparse graptolite and brachiopod fragments. Contains conodonts of the North Atlantic Province fauna high E to 2, of late Ibexian to early Whiterockian (Arenigian) age (Drake and others, 1988). These are deep-water clastic and carbonate rocks transported westward over the Laurentian carbonate shelf sequence during the Taconic orogeny. Most of the shale is strongly weathered to a depth of 10 feet or more. Thickness is approximately 1500 feet.

Lehigh Valley Sequence

Beekmantown Group, upper part - Dolomite, thinto thick-bedded, aphanitic to medium-grained, medium-light- to medium-gray, weathering lightto medium-gray or yellowish-gray; locally laminated and slightly fetid. Grades downward into medium- to thick-bedded dolomite, mediumto coarse-grained, medium-dark- to dark-gray, strongly fetid; has a mottled weathered surface and contains pods and lenses of dark-gray to black chert. Unit includes the Rickenbach Formation and Epler Formation of Drake and Lyttle (1985) and is the Ontelaunee Formation of Markewicz and Dalton (1977). Thickness ranges from 0-200 feet below the fault contact with rocks

of the Peapack klippe.

Beekmantown Group, lower part - interbedded dolomite and minor limestone, very thin- to thickbedded. Upper laminated, fine- to mediumgrained dolomite is very thin- to thick-bedded light-olive-gray to dark gray, and weathers darkyellowish-orange in places. Middle sequence is fine-grained dolomite with silty dolomite laminae, and thin- to medium-bedded, fine-grained limestone. Lower sequence consists of aphanitic to coarse-grained, thinly laminated to thickbedded, fetid dolomite with quartz sand laminae and local very thin to thin black chert beds. Corresponds to the Stonehenge Formation of Drake and Lyttle (1985) and to the Epler Formation and Rickenbach Formation of Markewicz and Dalton (1977). Thickness is approximately 250 feet in the map area.

Allentown Dolomite - Light- to dark-gray rhythmically bedded dolomitic mudstone, oolitic grainstone, algal laminate, and stromatolitic dolomite. Mudcrack brecciation common in upper part of cycles (Markewicz and Dalton 1980). Weathers yellowish-tan to brownish-gray, commonly in alternating layers 2 to 8 inches thick. Red hematite stain common in dolomite adjacent to the Mesozoic unconformity. Thickness is about 400 feet in map area.

Leithsville Formation - Interbedded light- to

medium-gray fine- to coarse-grained dolomitic limestone and calcitic limestone, thin-bedded light-gray to tan phyllite, and thin beds of dolomitic sandstone. Upper part (about 350 feet thick) is mostly fine- to medium-grained crystalline dolomite with irregular, intraclastic, and brec ciated texture; weathers dark gray, commonly with nodular or lumpy surface. Middle part (about 100 feet thick in map area) is light-tan to mediumgray silty to sandy dolomite, calcareous phyllite, purple to green siliceous or cherty dolomite, thin-bedded shale, quartzite, and dolomite Small-scale folds are common in shaly beds of the middle Leithsville Formation. Lower part (about 80 to 90 feet thick) consists mostly of dark to medium-gray, thick-bedded, medium- to finegrained dolomitic limestone with abundant stylolites and calcite-filled brecciation cracks. Large-scale solution features, including sinkholes and bedrock pinnacles are common in places near the top of the lower part, 70 to 100 feet stratigraphically above base of the formation. Contact with underlying Hardyston Quartzite is gradational. Thickness of the formation is about 550 feet.

Hardyston Quartzite - Pinkish-tan to white, medium- to very coarse-grained quartzose sandstone, conglomeratic in part, with 10 to 20 percent feldspar, mostly microcline (Aaron, 1969). Framework grains are moderately well sorted, mostly subangular (quartz) to angular (feldspar), and tightly cemented with silica Weathers tan to yellowish-brown, with blocky parting and fracture. Weathered surfaces are usually pitted due to dissolution of feldspar grains. Highly resistant to erosion; forms a conspicuous topographic bench in many areas with $10\,to\,20\,feet$ of relief . Thickness is about $40\,to\,60$ feet in map area.

Middle Proterozoic Rocks of the Reading Prong

Byram Intrusive Suite

Hornblende granite - Medium- to coarse-grained pink to buff, gneissoid to indistinctly foliated granite and sparse granite gneiss composed principally of microcline microperthite, quartz, oligoclase, and hornblende. Locally includes small bodies of pegmatite and amphibolite not shown on map.

Microperthite alaskite - Medium- to coarsegrained, pink to buff, gneissoid to indistinctly foliated granite composed principally of microcline microperthite, quartz, and oligoclase. Locally includes bodies of amphibolite not shown on map.

Metasedimentary Rocks

Potassic feldspar gneiss - Medium fine- to medium-grained, light pinkish-gray or buff, indistinctly foliated gneiss containing quartz, potassic feldspar, oligoclase, minor biotite, and opaque minerals. Contains thin amphibolite layers in some places.

Pyroxene gneiss - Medium-fine- to mediumgrained, greenish-gray, white- to tan-weathering, well-layered gneiss composed of oligoclase, clinopyroxene, variable amounts of quartz, and trace amounts of opaque minerals and sphene.

Biotite-quartz-feldspar gneiss - Medium-grained, gray to tan, locally rusty-weathering, well-layered and foliated gneiss, variable in texture and composition. Composed of oligoclase, quartz, biotite, and lesser microcline microperthite. Locally contains garnet, graphite, and sillimanite.

Amphibolite layers are common. **Losee Metamorphic Suite**

Quartz-oligoclase gneiss - Medium- to coarsegrained, white to light greenish-gray, poorly foliated granofels to moderately layered and foliated gneiss composed of quartz, oligoclase or andesine, and sparse to moderate amounts of hornblende and/or clinopyroxene. Locally contains amphibolite layers.

Biotite-quartz-oligoclase gneiss - Medium fine- to medium coarse-grained, light- to medium-gray or light greenish-gray, massive, moderately layered and foliated gneiss containing oligoclase or andesine, quartz, biotite, and locally garnet. Commonly interlayered with amphibolite.

Rocks of Uncertain Origin

Diorite - Medium- to coarse-grained, greenish-gray to brownish-gray, greasy-lustered, massive to moderately layered and foliated diorite to quartz diorite composed of andesine or oligoclase, clinopyroxene, hornblende, hypersthene, sparse biotite, and variable amounts of quartz. Amphibolite and mafic-rich quartz-plagioclase gneiss layers are common.

Amphibolite - Medium-grained, gray to greenishblack gneiss composed of hornblende and andesine. Locally contains biotite and/or clinopyroxene. Associated with almost all units of Middle Proterozoic age in the map area.

Pottersville Mylonite - Heterogeneous sequence of gray, light-green, or whitish-tan altered and retrogressively metamorphosed gneiss with mylonitic to ultramylonitic fabric. Interpreted to have been mafic-poor and mafic-rich varieties of quartz-plagioclase gneiss, amphibolite, and lesser hornblende granite and microperthite Rocks of Uncertain Age

> Diabase dike - Dense, massive, dark greenish-gray, fine-grained, altered and retrogressively metamorphosed diabase. Radiometric age has not been determined and could reasonably be between Middle Proterozoic and early Mesozoic.

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MAP SYMBOLS

—— — Contact - Dashed where inferred

Fault - Dashed where inferred, dotted where concealed, queried where uncertain

Strike-slip fault - Arrows indicate relative horizontal

displacement ligh-angle fault - U indicates upthrown side, D,

Thrust fault - Saw teeth on upper plate

Shear zone

Synform - Showing axis and direction of plunge

Antiform - Showing axis and direction of plunge Overturned synform - Showing axis, direction and

Overturned antiform - Showing axis, direction and dip of limbs, and direction of plunge

Strike and dip of crystallization foliation

dip of limbs, and direction of plunge

Strike and dip of mylonitic foliation

Strike and dip of bedding

Horizontal

Strike and dip of joints

Strike and dip of spaced cleavage

Strike and dip of slaty cleavage

Proterozoic rocks

Bearing and plunge of mineral lineation in

Bearing and plunge of intersection of

bedding and slaty cleavage

Abandoned magnetite mine

Abandoned stone quarry

SELECTED REFERENCES

Aaron, J.M., 1969, Petrology and origin of the Hardyston Quartzite (Lower Cambrian) in eastern Pennsylvania and western New Jersey, in Subitzky, Seymour, ed., Geology of selected areas in New Jersey and eastern Pennsylvania and guide to excursions: New Brunswick, N.J., Rutgers University Press, p. 21-34.

Bayley, W.S., Salisbury, R.D., and Kummel, H.B., 1914, Raritan, New Jersey: U. S. Geological Survey Geologic Atlas, Folio 191, 38 p., 5 maps at scale 1:125,000.

Drake, A.A., Jr., and Lyttle, P.T., 1985, Geologic map of the Blairstown quadrangle, Warren County, New Jersey: U. S. Geological Survey Geologic Quadrangle Map GQ-1585, scale 1:24,000.

Drake, A.A., Jr., Sinha, A.K., Laird, J., and Guy, R.E., 1988, The Taconic orogen, in Hatcher, R.D., Jr., Thomas, W.A., and Viele, G.W., eds., The Appalachian-Ouachita orogen in the United States: Boulder, Co., Geological Society of America, The Geology of North America, v. F-2.

Ethington, R.L., Furnish, W.M., and Markewicz, F.J., 1958. Ordovician conodonts in New Jersey: Journal of Paleontology, v. 32, p. 763-765.

Fedosh, M.S., and Smoot, J.P., 1988, A cored stratigraphic section through the northern Newark basin, New Jersey, in Froelich, A.J., and Robinson, G.R., Jr., eds., Studies of the Early mesozoic basins of the eastern United States: U.S. Geological Survey Bulletin 1776, p.19-23.

Kummel, H.B., 1896, Unpublished field notes: N. J. Geological Survey file maps, scale 1:21,120.

Lyttle, P.T., and Epstein, J.B., 1987, Geologic map of the Newark

1° x 2° quadrangle, New Jersey, Pennsylvania, and New York: U.S. Geological Survey Miscellaneous Investigations Series I-1715, scale 1:250,000.

Manspeizer, Warren, and Olsen, P.E., 1981, Rift basins of the passive margin: Tectonics, organic-rich lacustrine sediments, basin analysis, in Hobbs, G.W., ed., Field guide to the Paleozoic, Mesozoic, and Tertiary rocks of New Jersey and the central Hudson Valley: New York, Petroleum Exploration Society of New York, p. 25-105.

Markewicz, F.J., and Dalton, Richard, 1977, Stratigraphy and applied geology of the lower Paleozoic carbonates in northwestern New Jersey: 42nd Annual Field Conference of Pennsylvania Geologists Guidebook, 117 p.

Markewicz, F.J., and Dalton, Richard, 1980, Lower Paleozoic carbonates: Great Valley, in Manspeizer, Warren, ed., Field studies of New Jersey geology and guide to field trips: N.Y. State Geological Association 52nd Annual Meeting, Newark. N.J., Rutgers University, p. 54-69.

Minard, J.P., 1959, The geology of the Peapack-Ralston Valley in north-central New Jersey (unpublished M.S. thesis): New Brunswick, N.J., Rutgers University, 103 p.

Drake, A.A., Jr., and Berry, W.B.N., 1979, The Taconides of western New Jersey: New evidence from the Jutland klippe: Geological Society of America Bulletin, part 2, v. 90,

Perissoratis, Constantine, Brock, P.W.G., Brueckner, H.K.

ELEVATION ELEVATION JTapst **J**Ticsh **J**Tkpm MEAN SEA LEVEL 1000 Ojtb Yba MEAN SEA LEVEL No vertical exaggeration. Form lines indicate foliation in Proterozoic rock and bedding in Mesozoic rock. No vertical exaggeration. Form lines indicate foliation in Proterozoic rock and bedding in Mesozoic rock