

DESCRIPTION OF MAP UNITS

- af** ARTIFICIAL FILL—Excavated till, sand, gravel, and rock; and construction debris, cinders, and slag. In highway and railroad embankments, dams, and filled land. As much as 40 feet thick, generally less than 20 feet thick. Small areas of fill in urban areas are not mapped.
- Postglacial Deposits**—Sediment deposited along streams, in wetlands, and at the base of cliffs. These deposits have been accumulating since retreat of the late Wisconsinan glacier.
- Qal** ALLUVIUM—Silt, clay, sand, and pebble-to-cobble gravel. Contains variable amounts of organic matter. Color of fine sediment is gray, brown, and yellowish brown. Fine sediment was deposited as overbank material on the floodplains and may be as much as 15 feet thick along major streams. It generally overlies gravel deposited in the stream channel. The gravel is generally less than 5 feet thick. In the valleys of Black, Fochtick, and Wawayanda Creeks, alluvium may overlie swamp and marsh deposits.
- Qs** SWAMP AND MARSH DEPOSITS—Gray silt and clay overlain by dark-brown peat. May include some fine-grained alluvium along larger streams. As much as 25 feet thick (Wakeman and others, 1943).
- Qta** TALUS—Angular boulders and blocks of gneiss, with little or no matrix material, forming a steep apron at the base of cliffs on the west side of Wawayanda Mountain. Maximum thickness 20 feet (estimated).
- Qst** STREAM TERRACE DEPOSITS—Sand, silt, and pebble gravel, minor cobble gravel, forming terraces 5 to 10 feet above the present floodplain. Fine sediment is brown to yellowish brown. Maximum thickness 15 feet (estimated).
- Ql** ALLUVIAL FAN DEPOSITS—Cobble-to-boulder gravel, sand, silt, and fine sediment is grayish brown, brown, and yellowish brown. Form fan-shaped deposits at the mouths of canyons and gullies eroded into thick glacial sediment. As much as 30 feet thick (estimated).
- Glacial Deposits**—Glacial deposits in the map area include till and stratified sand, gravel, silt, and clay. The till occurs in drumlins, moraines, and as a layer of variable thickness on the bedrock surface. It was deposited directly from glacial ice and is as much as 120 feet thick. The stratified sediment occurs discontinuously as knolls, ridges, and plains in valleys and lowlands. It was deposited by glacial meltwater in glacial lakes and river plains and is as much as 120 feet thick.
- These sediments were deposited by the late Wisconsinan glacier. As indicated by the orientation of striations and drumlins, and the distribution of erratics, this glacier advanced southerly to southwesterly across the map area. It reached its maximum position about 20 miles south of the southern edge of the map area. During advance the glacier deposited till in drumlins and locally on hillsides that faced the advancing ice; elsewhere, it eroded bedrock. The erosional features include scoured troughs in weak carbonate and shale bedrock in the vicinity of Black Creek and Upper Greenwood Lake, and extensive areas of abraded ledges (on slopes that faced toward advancing ice) and quarried cliffs (on slopes that faced away from advancing ice) on resistant gneiss and quartzite bedrock on Pechuck, Wawayanda, and Bearfoot Mountains.
- Late Wisconsinan ice began to melt back from its terminal position approximately 20,000 years ago. As the ice front retreated northeastward, glacial lakes formed in valleys that sloped toward the glacier, and were thus dammed, or in valleys that were dammed by previously-deposited glacial sediment. Valleys that were not dammed carried rivers of glacial meltwater. In a few places till and related nonstratified sediment were deposited in small moraines along the ice front. The position and elevation of stratified deposits and lake spillways, and the trend of moraines, permit an approximate determination of the orientation of recessional ice margins. In the map area, the trend of the Mud Pond and Cherry Ridge moraines indicates generally east-west-trending recessional ice margins. Topographically-induced lobation was slight on Wawayanda Mountain, but probably more pronounced along the east and west walls of Vernon Valley. Details of the glacial lake elevations and spillways are provided in the map unit descriptions.
- Deposits in Glacial Lakes**—These deposits are stratified and generally well sorted. They include sand and gravel deposited in deltas and lacustrine fans, and clay, silt, and fine sand deposited on lake-bottom plains.
- GLACIAL LAKE WALKKILL DEPOSITS**—Deltic (Qw4), lacustrine-fan (Qw4), and lake-bottom (Qw4) sediment deposited in glacial Lake Walkkill. Lake Walkkill occupied the north-draining Walkkill basin. It was controlled by a spillway at an elevation of 500 feet at the low point on the Delaware-Walkkill divide near Augusta, New Jersey (14 miles southwest of Vernon). It lowered when three successively lower spillways formed on the Walkkill-Hudson divide, declining in elevation from 440 to 370 to 350 feet, were uncovered by the retreating ice front at and north of Goshen, New York, about 20 miles north of Vernon. The 440-foot level may have extended into the map area as a shallow lake in the lowest parts of Vernon Valley; the lower levels did not. As the ice from retreated farther north, Lake Walkkill fell to even lower levels (Connally and Sifkin, 1967).
- Qw4** Sand, pebble-to-cobble gravel, minor silt. As much as 120 feet thick.
- Qw3** Sand, pebble-to-cobble gravel, minor cobble-to-boulder gravel. As much as 50 feet thick.
- Qw1** Silt, clay, and fine sand. As much as 100 feet thick.
- WAWAYANDA DEPOSITS**—Deltic sediment deposited in six separate glacial lakes ponded in north-draining valleys on Wawayanda Mountain. Lake-bottom deposits are not exposed but are probably present beneath the extensive swamp and marsh deposits (unit Qm) covering the bottoms of the former lake basins.
- Qw5** Pebble-to-cobble gravel and sand. As much as 40 feet thick (estimated). Spillway drained outward into Lake Walkkill at an elevation of about 770 to 800 feet.
- Qw5** Pebble gravel and sand, minor fine cobble gravel. As much as 20 feet thick (estimated). Spillway drained outward into glacial Lake Bearfoot at an elevation of 1130 feet.
- Qw4** Pebble-to-cobble gravel and sand. As much as 50 feet thick. Spillway drained outward into glacial Lake Bearfoot at an elevation of 1150 feet.
- Qw3** Pebble-to-cobble gravel and sand. Maximum thickness 30 feet. Spillway drained outward into glacial Lake Bearfoot at an elevation of 1170 feet.
- Qw2** Pebble gravel and sand. Maximum thickness 60 feet (estimated). Spillway drained outward into glacial Lake Bearfoot at an elevation of 1230 feet.
- Qw1** Pebble gravel and sand. Maximum thickness 60 feet (estimated). Spillway drained outward into the Mossman Brook valley at an elevation of 1250 feet.
- GLACIAL LAKE BEARFOOT DEPOSITS**—Deltic (Qb1, Qb2) and lake-bottom (Qb1) sediment deposited in glacial Lake Bearfoot. Lake Bearfoot occupied the north-draining valley of Longhouse Creek. It was controlled by an early spillway to the south draining into Mossman Brook (the Qb1 deposits), and, later, by a lower spillway draining outward into the Wawayanda basin (the Qb2 deposits). The lake lowered and drained as lower spillways draining westward into the Walkkill basin were uncovered by the retreating ice margin in New York state, north of the map area.
- Qb2** Pebble-to-cobble gravel, minor boulder gravel, overlying sand. As much as 120 feet thick. Spillway drained outward into the Wawayanda valley at an elevation of 1090 feet.
- Qb1** Pebble-to-cobble gravel and boulder gravel overlying sand. As much as 80 feet thick. Spillway drained outward into the Mossman Brook valley at an elevation of 1140 feet.
- Qb1** Fine sand, silt, minor clay. As much as 50 feet thick. Deposited in both stages of Lake Bearfoot; generally underlies deltaic sand and gravel.
- Qs** UNNAMED DEPOSIT—Sand and pebble gravel in small ponded valley 2 miles south of Vernon. Maximum thickness 15 feet.
- Qmp** MARSHALL POND DEPOSIT—Pebble-to-cobble gravel and sand, minor boulder gravel. Maximum thickness 20 feet (estimated). Spillway drained to south at an elevation of approximately 1220 feet, either over stagnant ice or moraine deposits of unit Qmp.
- Qv3** UNION VALLEY DEPOSIT—Deltic sand and pebble gravel. Maximum thickness in map area 20 feet (estimated). Deposited in a glacial lake occupying the north-draining valley of Belcher Creek. Controlled by a spillway on the Wawayanda-Pequanock divide at an elevation of 850 feet near Postville, approximately 2 miles southwest of Pinecliff Lake (Stanford, 1991). Qv1 and Qv2 are restricted to the Newfoundland quadrangle.
- Deposits of Glacial Streams**—Stratified, generally well sorted gravel and sand forming valley-bottom plains, terraces, and fans in valleys not occupied by glacial lakes.
- Qm1** MELT-WATER FAN DEPOSITS—Cobble-to-boulder gravel and sand in fan-shaped deposits at mouths of meltwater channels. Maximum thickness 30 feet (estimated).
- Qcl** CLINTON BROOK OUTWASH—Boulder-to-cobble gravel. Maximum thickness 20 feet (estimated).
- Qpc** PACACK BROOK OUTWASH—Boulder-to-cobble gravel. Maximum thickness 20 feet (estimated).

PEQUANNOCK OUTWASH—Boulder-to-cobble gravel. Maximum thickness 20 feet (estimated).

- Qps** KITTATINNY TILL—Light-brownish-gray to very-pale brown silt to sandy silt with some to many subrounded to subangular pebbles and cobbles and few to some subrounded boulders. Gravel includes gray carbonate rock and gray mudstone and sandstone, with minor white-to-gray quartzite and quartzite conglomerate and a trace of gneiss. The gneiss becomes more abundant near contact with unit Qs. The carbonate, mudstone, and sandstone clasts are derived from Paleozoic sedimentary rocks of the Walkkill valley. The quartzite and quartzite conglomerate are derived from the Shawangunk Formation to the north and west of the Walkkill valley, and to a lesser extent from local outcrops of the Hardyston Formation. Boulders are generally quartzite and carbonate rock. Unit Qs is as much as 120 feet thick but is generally between 20 and 50 feet thick. Unit Qs delineates areas where the till is discontinuous and generally less than 20 feet thick over bedrock.
- Qn** NETCONEG TILL—Yellow to yellowish-brown silt and sand with many subrounded to subangular pebbles and cobbles and some to many subrounded boulders. Gravel includes chiefly gneiss and gray mudstone and sandstone, with minor gray carbonate rock and white-to-gray quartzite. Gneiss is the local bedrock; the other clasts are derived from bedrock in the Walkkill valley north of Shawangunk Mountain to the north of the map area. Boulders are chiefly gneiss, with a few scattered carbonate boulders, some of which are shown by a special symbol on the map. East of the Mossman Brook-Longhouse Creek valley, the till matrix is somewhat silty and roller; here, gray and red shale and siltstone are abundant in the gravel fraction. On Bearfoot Mountain, purple quartzite and quartzite-conglomerate are abundant in both the gravel and boulder fractions. These differences reflect incorporation of material derived from the underlying Bellevue Sandstone and Skumunk Conglomerate, which crop out in the valleys and on Bearfoot Mountain (Herman and Mitchell, 1991). Unit Qn is as much as 50 feet thick, but is generally 20 to 50 feet thick. Unit Qn delineates areas where the till is discontinuous and generally less than 20 feet thick over bedrock.
- Qncr** CHERRY RIDGE MORaine—Netcong till, generally with a sandy matrix and containing many to very many subangular cobbles and boulders of gneiss. Forms a discontinuous linear belt of low hummocks, shallow basins, and acute ridges extending from Highland Lake eastward to Lookover Lake. The hummocky topography generally has a relief of less than 15 feet but relief may be as much as 30 feet. The acute ridges are 10 to 40 feet high. Several of them, shown by a special symbol on the map, are asymmetric in profile, with gentle north slopes and steeper south slopes. This suggests that they formed at an active ice front. Maximum thickness of till in the moraine is estimated to be 50 feet.
- Qmud** MUD POND MORaine—Netcong till as in unit Qncr, although in places, particularly just east of the Sussex-Passaic county line, the deposits are extremely bouldery and contains little or no matrix material. Forms a discontinuous belt of low hummocks, shallow basins, and a few acute ridges extending from west of Canister Reservoir eastward to Urtown. The hummocky topography generally has a relief of less than 15 feet but relief may be as much as 40 feet. The ridges are 10 to 20 feet high. One small ridge near Urtown has an asymmetric profile as described in unit Qncr. Maximum thickness of till in the moraine is estimated to be 50 feet.
- ICE-CONTACT DEPOSITS**—Three small deposits of nonstratified, nonconformable sand and silt with much subangular gravel, locally bouldery. Forms low ridges or knolls less than 20 feet tall. These deposits may be sediment flows from the glacial surface deposited during glacial recession.
- Qic** Contact—Dashed where gradational or feathering, dotted where concealed or excavated.
- Striation**—Location at dot. Where two sets of striae occur the earlier set is indicated by a broken arrow.
- Meltwater channel**—Narrow, linear, bouldery drainage way in which present streams are nonexistent or underflow.
- Scarp cut by meltwater**—Line at top of scarp, ticks on slope.
- Scarp cut by postglacial streams**—Line at top of scarp, blocks on slope.
- Artificial excavation scarp**—Line at top of slope, ticks on slope.
- Crest of asymmetric ice-contact ridge**—Barbs on gentle, concave slope.
- Esker**
- Drumlin**
- Erratic boulder of carbonate rock**—Indicates southerly glacial transport from outcrops of carbonate bedrock in Vernon Valley and in the Walkkill valley north of the map area. Not all occurrences shown.
- Spillway for glacial lake**—Lettering indicates associated deposit.
- Gravel pit—Active in 1990.**
- Gravel pit—Inactive in 1990.**
- Elevation of the bedrock surface**—Contour interval 50 feet, 100 feet where data are sparse. Shown only in major valleys. Hashmarks indicate closed contours in glacially-scoured basins.
- Well on section**—Number refers to well number in table 1.
- Well with log in table 1**
- Area of extensive bedrock outcrop—surficial sediment generally absent.**
- Bodies of water**—Not shown on base map.
- Unit on left is thin and overlies unit on right**
- Both map units are alternately present and thin**
- Surface accumulation of boulders—Till surfaces washed by subglacial, proglacial, or ice-marginal meltwaters. Does not include talus or bouldery moraines.**

REFERENCES

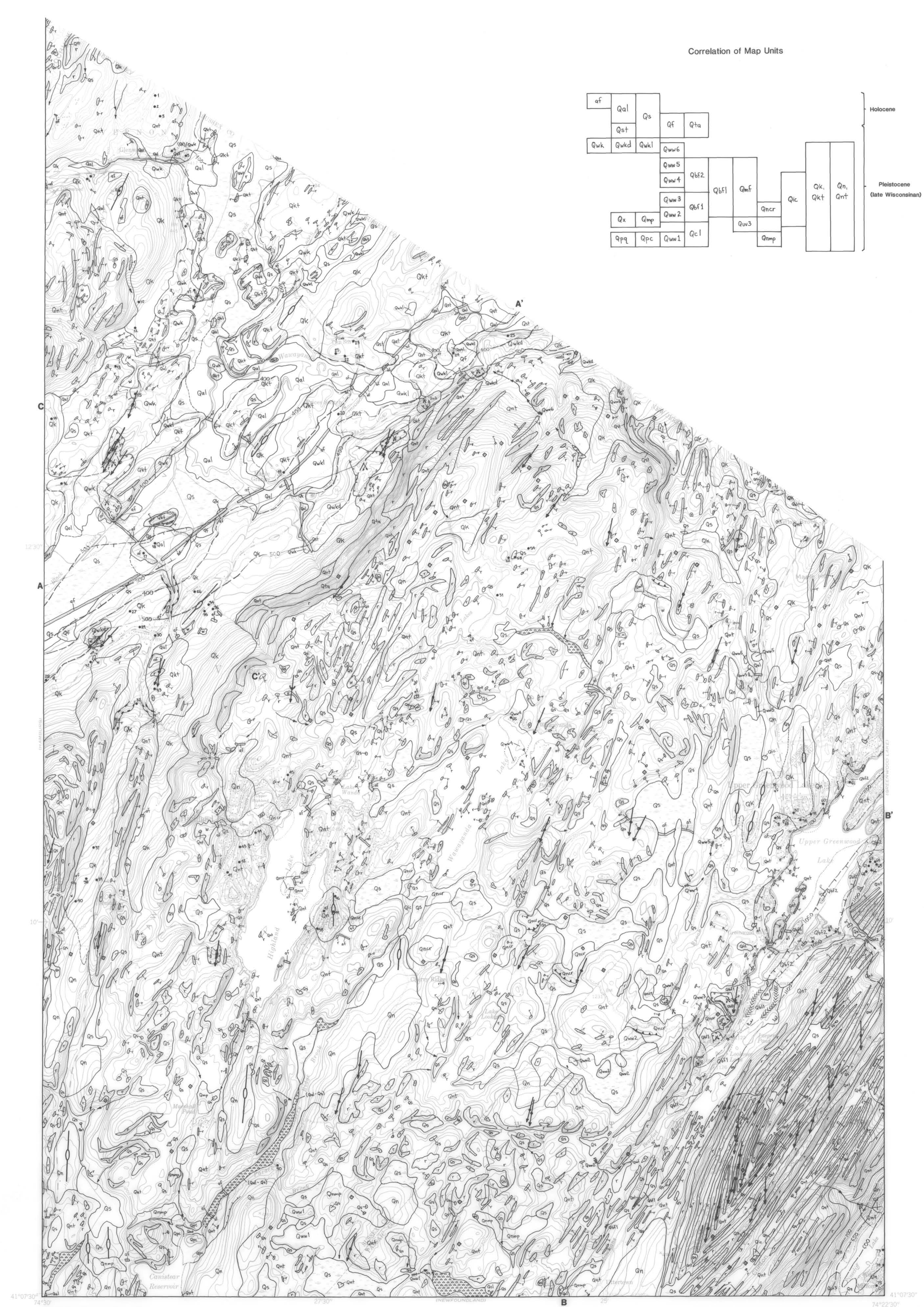
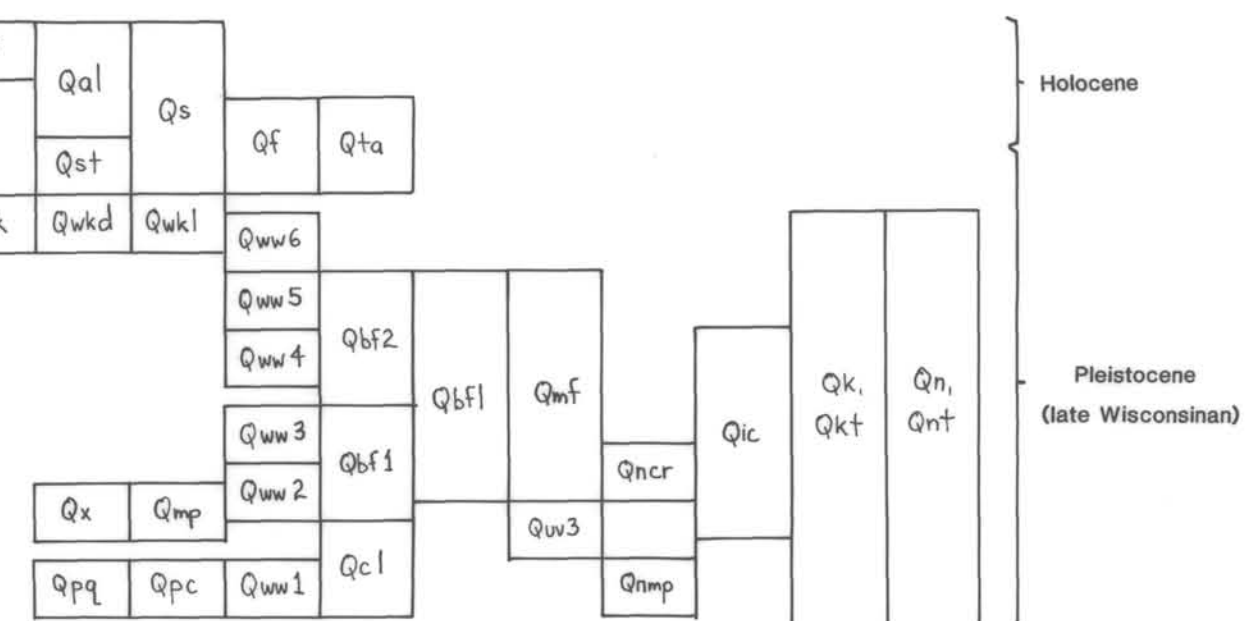
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- Wakeman, S. A., Schellhoff, H., Hickman, C. A., Carlson, T. C., and Stevens, S. C., 1943, The peats of New Jersey and their utilization: N. J. Department of Conservation and Development Geologic Series, Bulletin 55, Part B, 278 p.

Table 1—Selected Well Logs	
Well No.	Depth (feet)
1 22-20384	0-30
2 22-20803	0-48
3 22-19423	0-70
4 22-19128	0-25
5 22-26153	0-57
6 22-21553	0-79
7 22-18881	0-50
8 22-13363	0-71
9 22-5881	not reported
10 22-18930	0-50
11 22-18068	0-50
12 22-18696	0-47
13 22-18554	0-50
14 22-18793	0-108
15 22-19589	0-30
16 22-17863	0-75
17 22-21151	0-70
18 22-17426	0-5
19 22-21837	0-38
20 22-19046	0-40
21 22-20177	0-61
22 22-22667	0-2
23 22-19467	0-14
24 on file at N.J. Geological Survey	0-45
25 22-22685	0-112
26 22-18584	0-110
27 22-2907	0-115
28 22-4287	0-30
30 on file at N.J. Geological Survey	0-16
31 22-20607	0-26
32 22-4740	0-46
33 22-6696	0-40
34 22-19131	0-20
35 22-22071	0-55
36 22-20534	0-46
37 22-19562	0-110
38 22-22621	0-55
39 22-23664	0-70
40 22-17990	0-20
41 22-1471	0-51
42 22-1420	0-62
43 22-20540	0-32
44 22-18092	0-50
45 22-18094	0-50
46 22-20692	0-27
47 22-20544	0-28
48 22-18020	0-26
49 22-22496	0-53
50 22-20559	0-21
51 22-20356	0-32
52 22-19878	0-20
53 22-21132	0-15
54 22-19968	0-30
55 22-16460	0-45
56 22-2965	0-10
57 22-538	0-76
58 22-460	0-72
59 22-626	0-85
60 22-3148	0-65
61 22-22912	0-15
62 22-1294	0-80
63 22-2329	0-72
64 22-2973	0-42
65 22-2206	0-40
66 22-436	0-81
67 22-4114	0-81
68 22-1743	0-68
69 22-18481	0-2
70 22-19271	0-20
71 22-2125	0-68
72 on file at N.J. Geological Survey	0-180
73 22-10149	0-71
74 22-1860	0-30
75 22-22673	0-140

(1.) Well permit issued by the New Jersey Department of Environmental Protection and Energy, Bureau of Water Allocation.

(2.) Inferred map units and comments in parentheses.

Correlation of Map Units



Base from U.S. Geological Survey, 1954

SCALE 1:24,000  
1000 2000 3000 4000 5000 6000 7000 8000 9000 10000 FEET  
1:24,000  
1000 2000 3000 4000 5000 6000 7000 8000 9000 10000 METERS

CONTOUR INTERVAL 20 FEET  
DATUM IS MEAN SEA LEVEL

1918 GRID AND 1983 MAGNETIC NORTH  
DECLINATION AT CENTER OF SHEET

QUADRANGLE LOCATION

Geology mapped in 1987, 1989-91

The interpretations presented here are provisional pending  
re-evaluation. There may be revisions prior to publication.

**SURFICIAL GEOLOGY OF THE  
WAWAYANDA AND PINE ISLAND QUADRANGLES,  
SUSSEX AND PASSAIC COUNTIES, NEW JERSEY**  
by  
Scott D. Stanford  
1992

